

## Groundwater / Porous Flow Review

Aquifer - a geologic formation that (1) contains and (2) transmits water (2gpm) under normal field conditions

Aquiferard - same, but transmits at slow rate as compared to aquifer

Aquiclude - cannot transmit water

Aquifuge - cannot contain water

Porosity - ratio of void volume to aquifer volume  $n = \frac{V_{\text{void}}}{V_{\text{bulk}}}$

Water content -  $\theta_w = \frac{V_{\text{water}}}{V_{\text{bulk}}}$   $\theta_w = n$  if completely saturated

Confined aquifer - bounded above & below by relatively impervious formations. Formations are upper/lower flow boundaries

Unconfined aquifer - upper flow boundary is free surface

Leaky aquifer - gains/loses water to adjacent aquifers through vertical leakage

Hydraulic conductivity - measure of relative ease of flow through an aquifer. Combined property of porous medium and fluid,  $K$

Transmissivity - measure of aquifer's ability to transmit water through its entire thickness

Storage - amount of water added to/released from storage per unit change in head per unit area of aquifer  

$$S = \frac{V_w}{A\Delta h}$$

Mechanisms - confined: aquifer & water compression/decompression  
 unconfined: same + drainage/filling pore space

Confined:  $S$  - storage coefficient

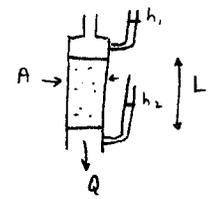
Unconfined:  $S_y$  - specific yield  
 $S_r$  - specific retention }  $n = S_y + S_r$

Darcy's Law (Equation of Motion)

$$Q = -KA \left( \frac{h_2 - h_1}{L} \right)$$

$$\lim_{L \rightarrow 0} \frac{h_2 - h_1}{L} = \frac{dh}{dL}$$

$$\therefore Q = -KA \frac{dh}{dL}$$



Specific discharge is  $U = \frac{Q}{A}$ . Average linear velocity is  $U = \frac{U}{n} = \frac{Q}{nA}$

In 3-D  $\underline{Q} = (Q_x, Q_y, Q_z) = -\underline{K}A \cdot \text{grad}(h) = -A \underline{K} \cdot \nabla h$

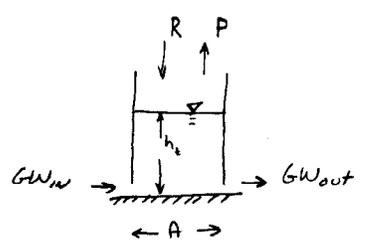
$$\underline{U} = (U, V, W) = -\underline{K} \cdot \nabla h$$

$$\underline{V} = (v, v, w) = \frac{1}{n} \underline{U} = \left( \frac{U}{n}, \frac{V}{n}, \frac{W}{n} \right) = -\frac{1}{n} \underline{K} \cdot \nabla h$$

Cell Balance Approach

Treats aquifer as linked single cell models.

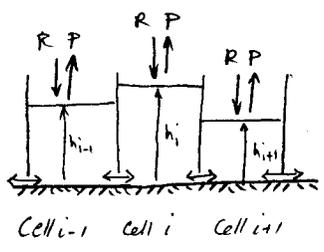
- Write cell balance for each cell
- Use Darcy's law to link the cells
- Generalize expressions
- Apply solution algorithm to resulting equations



Generic cell

$$SA \frac{dh}{dt} = AR - P + GW_{in} - GW_{out}$$

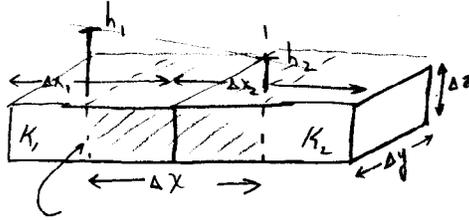
Linked cells



$$S_i A_i \frac{dh_i}{dt} = A_i R_i - P_i + GW_{i-1 \rightarrow i} - GW_{i \rightarrow i+1}$$

Groundwater flow terms

$$Q_{1 \rightarrow 2} = \bar{K} \Delta y \Delta z \frac{h_1 - h_2}{\Delta x}$$



Apply Darcy's law in shaded area

$$\bar{K} = \frac{K_1 + K_2}{2} \text{ (Arithmetic mean, OK if } K_1 \approx K_2 \text{)}$$

$$\bar{K} = \frac{\Delta x_1 + \Delta x_2}{\frac{\Delta x_1}{K_1} + \frac{\Delta x_2}{K_2}} \text{ (Harmonic mean, Use if cells are big and } K_1 \neq K_2 \text{)}$$

$$Q_{1 \rightarrow 2} = \underbrace{\bar{K} \Delta z}_{\bar{T}} \frac{\Delta y}{\Delta x} h_1 - h_2 \quad \text{but } \Delta x = \Delta x_1 + \Delta x_2$$

if  $\Delta x_1 = \Delta x_2$  then

$$Q_{1 \rightarrow 2} = \bar{T} \frac{\Delta y}{\Delta x} (h_1 - h_2) = - \bar{T} \frac{\Delta y}{\Delta x} \Delta h \quad \Delta h = h_2 - h_1$$

$$\therefore S_i A_i \frac{dh_i}{dt} = A_i R_i - P_i + \bar{T}_{i-1} \frac{\Delta y}{\Delta x} (h_{i-1} - h_i) - \bar{T}_{i+1} \frac{\Delta y}{\Delta x} (h_i - h_{i+1})$$

 $A_i = \Delta x_i \Delta y_i$  divide by  $A_i$ , use arithmetic mean as example

$$S_i \frac{dh_i}{dt} = R_i - \frac{P_i}{\Delta x_i \Delta y_i} + \left( \frac{T_{i-1} + T_i}{2} \right) \left( \frac{h_{i-1} - h_i}{\Delta x^2} \right) - \left( \frac{T_i + T_{i+1}}{2} \right) \left( \frac{h_i - h_{i+1}}{\Delta x^2} \right)$$

In limit as  $\Delta x, \Delta y \rightarrow 0$ 

we have

$$S \frac{\partial h}{\partial t} = r - p - \text{div}(\underline{T} \cdot \text{grad}(h)) = r - p - \nabla \cdot (\underline{T} \cdot \nabla h)$$

(Same result using fluid mechanics principles)

Return to the cell balance equation - replace  $\frac{\partial h}{\partial t}$  by finite-difference approximation

$$S_i \frac{h_i^{\text{trst}} - h_i^t}{\Delta t} = R_i - \frac{P_i}{\Delta x_i \Delta y_i} + \left( \frac{T_{i-1} + T_i}{2} \right) \left( \frac{h_{i-1}^{\circ} - h_i^{\circ}}{\Delta x^2} \right) - \left( \frac{T_i + T_{i+1}}{2} \right) \left( \frac{h_i^{\circ} - h_{i+1}^{\circ}}{\Delta x^2} \right)$$

time level, determines  
implicit/explicit  
scheme